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At the time of the study, the mean age of the participants was 16.4 years (range 15.5–17.5 years).

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Banded iron rocks (BIFs) are the most important marine sedimentary rocks deposited during the Proterozoic. They consist of alternating layers of iron-rich (up to 15% iron) and iron-poor (less than 1% iron) minerals, typically chert (or its metastable polymorph, microcrystalline quartz) and iron minerals with the iron content of 10–30%. The first BIFs precipitated at 3.8 Ga (Isua, Greenland).

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models have been proposed for interpreting the formation of the Neoproterozoic BIFs: (1) During the break-up of Rodinia supercontinent

induced the addition of Fe flux from submarine exhalative (Basta et al., 2011; Cox et al., 2013); (2) The thick ice sheet formed by "Snowball Earth" event isolated the hydrosphere and oxidized atmosphere, which led to the reduction of ocean and the dissolution of a large amount of Fe element. After the ice cover melted, the Fe was oxidized and precipitated into BIF as a result of the contact between ocean and atmosphere (Hoffman et al., 1998; Hoffman and Schrag, 2002; Klein and Ladeira, 2004; Halverson et al., 2011) and (3) during the late Neoproterozoic, there was a reduced S-poor and Fe-rich environment, which was favorable for Fe element existing as divalent ions in seawater, and then to be oxidized and precipitated (Johnston et al., 2010).

In this study, we report a newly identified Neoproterozoic Baijianshan BIF at the southeastern margin of the Tarim Block in NW China. The precipitation mechanism and the oceanic redox condition of the Baijianshan BIF was investigated based on detailed field observations coupled with systematic elemental and isotopic compositions analysis. This study sheds a new light on our understanding of the possible coupling of the Cryogenian snowball Earth event and the formation of Neoproterozoic BIFs.

## **2 Regional Geology**

The Tarim Craton in NW China, covering an area of more than 600,000 km<sup>2</sup>, is one of the main three Precambrian nuclei in China (i.e., North China, South China and Tarim). It is surrounded by the Phanerozoic orogenic belts of Tianshan Mountains to the north, the western Kunlun Mountains to the south, and the Central-Southern Altyn Tagh Mountains to the southeast. Despite most of the Tarim Craton being covered by aeolian sands and difficulties in accessibility, previous work reveals that the craton is characterized by a typical double-layered structure consisting of a Pre-Cryogenian basement and Cryogenian-Cambrian cover sequences (Xinjiang BGMR, 1993; Zhang et al., 2013). Particularly, recent studies demonstrate that the craton docked at the northern fringe of the Rodinia slightly earlier than 760 Ma and then the Cryogenian-Cambrian sedimentary sequences deposited during the breakup of the Rodinia (Zhang et al., 2010, 2016; Xu et al., 2013).

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Titanium was analyzed using the procedure described by Li et al. (2000). The results were presented with 1 $\sigma$  and 2 $\sigma$  confidence limits (Ludwig, 2003). The zircon U-Pb data are listed in Table 1.

### Whole rock geochemistry

Forty-five BIF samples and three basalt samples were carefully selected along the profile. Whole-rock major compositions were analyzed using standard fused glass beads at the Nanjing Institute of Geology and Mineral Resources (NIGMR) following the procedures similar to those described by Li et al., (2000). The analytical errors between 1 and 5%. In addition, we used wet chemistry technique for Fe and Mn following the procedures described by Andrade et al. (2002). The mean values are listed in Table 2.

Trace elements were determined using a Perkin-Elmer Sciex ELAN DEQEX-2 ICP-MS. The detection limit Geochemistry, Institute of Geochemistry, Chinese Academy of Sciences. The analytical errors were better than 5%. Samples were digested in HF-HNO<sub>3</sub> in Teflon-lined stainless steel bombs at 190 °C for 24 h. The analytical errors were better than 1% with concentrations less than 100 ng/g. The results are reported in supplementary material.

Sr-Nd isotopes were analyzed using the Micromass TRIPLE-TRACE procedure at the Institute of Geochemistry, Chinese Academy of Sciences (Cheng et al., 2004). The analytical errors for mass-fractionation are less than 0.1%. The  $^{87}\text{Sr}/^{86}\text{Sr}$  and  $^{143}\text{Nd}/^{144}\text{Nd}$  ratios are reported as  $\epsilon_{\text{Nd}}$  values for  $t = 0$  Ma. The  $\epsilon_{\text{Nd}}$  values are calculated using the initial age of 4550 Ma for the Earth (DePaolo, 1988).

$$\delta^x\text{Fe}/^{56}\text{Fe}$$

The performance of the CAGS Fe (CAGS Fe) that yielded data for the measurements of an iron isotope reference material (Fe-1) and Fe-2.

$^{57}\text{Fe}_{\text{CAGS}}$

$^{57}\text{Fe}$

### 5.1 Zircon U-Pb age

Zircons from the tuff sample 2071 are variable in size with the length varying from 60 μm to 150 μm and the aspect ratios 1-2. In CL images most zircon exhibit oscillatory zoning, sharing the features of the zircons crystallized from silicic magma (Wu and Zheng, 2004). Thirty-two analyses were conducted on 32 zircon grains and the results are presented in supplementary Table 1 and illustrated in Fig. 4. Obviously, some analyses show variable radiogenetic lead lost. However, the results can be broadly divided into two sub-groups. Group 1, which is likely to be xenocrystal and generally has larger size, constructs a good Discordia with upper intercepted age of  $1378 \pm 44$  Ma (MSWD=0.92) (Fig. 4a). Group 2 mostly shows euhedral in form and smaller in size. Twenty-one analyses construct a Discordia with an intercepted age of  $736.2 \pm 3.8$  Ma ( $N=21$ , MSWD=1.08) (insert of Fig. 4a). The significant discordant analyses, the left fifteen analyses of well concordance between  $^{206}\text{Pb}/^{238}\text{U}$  and  $^{207}\text{Pb}/^{235}\text{U}$  (Fig. 4b), define a discordia with lower intercepted age of  $736.2 \pm 3.8$  Ma (MSWD=0.20). This age is interpreted as the crystallization age of the tuff.

### 5.2.1 Ge

Three samples have relatively low  $\text{TiO}_2$  contents (39.93% to 44.55%). The increase in  $\text{TiO}_2$  is caused due to the high LOI. Their high Nb/Y ratios (1.92 to 2.12) suggest that they are enriched in Nb/Y. In the Nb/Y vs. Zr/TiO<sub>2</sub> diagram, they plot into the field of Nb/Y enriched rocks. With respect to the trace elements, they have REE patterns ranging from REE depleted to REE enriched Chondrite-normalized patterns (Fig. 10). They have low Cr (214-306 ppm) and Ni contents (16-20 ppm). In the primitive mantle-normalized diagram, they show a significant Nb-Ta trough (Nb/La = 1.0-1.1) (Figure 10).

As shown in Table 2, the samples from Baijianshan BIF are rich in Fe<sub>2</sub>O<sub>3</sub> (13.11%–56.65% (38.44% on average)) and SiO<sub>2</sub> (35.06%–42.02% (42.02% on average)). They have low contents of Al<sub>2</sub>O<sub>3</sub> (0.03%–3.66% (on average)), TiO<sub>2</sub>, MnO, P<sub>2</sub>O<sub>5</sub>, CaO, Na<sub>2</sub>O, K<sub>2</sub>O and variable total rare earth elements (incompatible elements) (supplementary Table 2).

Their REE abundances vary from 38.1 ppm to 150.8 ppm. Normalized to the La<sub>N</sub>/Yb<sub>N</sub> of 0.23-0.69 (Fig. 5b). The REY patterns exhibit no of Ce/Ce\* ~ 1 and Pr/Pr\* ~ 1; Bau and Dulski, 1996) and in 0.96-1.14; Y/Ho ratios ranging from 23.3 to 31.8 with an av

Sr-Nd concentrations and isotopic ratios of the supplementary Table 3. As for the basalts, the basalts ( $^{87}\text{Sr}/^{86}\text{Sr}$ ); ranging from 0.7061 to 0.7084 du

6.9.

Strontium and neodymium concentrations range from 7.06 ppm to 59.8 ppm, respectively. The  $^{87}\text{Sr}/^{86}\text{Sr}$  ratios range from 0.51231–0.51243 and  $^{87}\text{Sr}/^{86}\text{Sr}$  (0.51231–0.51243) and  $^{87}\text{Sr}/^{86}\text{Sr}$  (0.51231–0.51243) to 1.46, with an average of 0.2. The  $^{87}\text{Sr}/^{86}\text{Sr}$  values are between 0.6412 and 0.6414, with a large range of Rb/Sr ratios.

The Fe isotope composition of the samples is presented in Supplementary Table 4. The six samples have  $\delta^{56}\text{Fe}$  values ranging from -0.1 to +0.2‰, with values of 1.78-1.82‰ for the two samples from the 1.8 Ga BIFs. Their enrichment in  $^{56}\text{Fe}$  is consistent with the values for Archean BIFs (Halverson et al., 2011; Yavuz et al., 2012) and is distinct from the values for Proterozoic BIFs (Halverson et al., 2011; Yavuz et al., 2012).

## 6 Diso

## What are BIFs

observations revealed that the Neoproterozoic Baijianshan BIF has not

experienced significant metamorphism (Fig. 3b, c and d), indicating the original geochemical characteristics were well retained and can be used to decipher the iron sources and the oceanic environments when it precipitated.

The Baijianshan BIF shows a modern seawater REE signature with significant LREE depletion relative to the HREE ( $\text{La}_N/\text{Yb}_N = 0.41$ ) (Fig. 5b, Zhang and Nozaki, 1996; Alibo and Nozaki, 1999). A range of factors can affect the primary REY composition in terms of BIF depositional processes, including post-depositional metamorphism and/or syn-deposition clastic contamination (Alexander et al., 2008; Hu et al., 2020). Previous studies demonstrated that diagenetic and metamorphic events (such as weathering and/or fluid-rock interactions) have negligible effects on the REY of the chemical sediments (Bolhar et al., 2004). However, Rb is a relatively mobile element compared to highly immobile elements such as Th. In Fig. 6a, the well positive correlation between Rb and Th demonstrates insignificant mobility of the LILEs (Large Ion Lithophile Elements) and the absence of significant diagenetic or metamorphic alteration effects, this conclusion is also consistent with thin section observations (Fig. 3c, d, g, h). With respect to the syn-depositional processes, the contents of  $\text{Al}_2\text{O}_3$  and  $\text{TiO}_2$  are consistently low with averages of 3.66% and 0.41%. In addition, despite the positive correlations between  $\text{Al}_2\text{O}_3$  and  $\text{TiO}_2$  ( $r = 0.93$ ),  $\text{K}_2\text{O}$  ( $r = 0.79$ ), Rb ( $r = 0.78$ ), V ( $r = 0.45$ ), Sc REE ( $r = 0.62$ ) (Fig. 6b-i) signify the incorporation of terrigenous contribution in the deposition processes, most correlation coefficients are much lower than those BIFs sourced from recycled crust (Basta et al., 2011; Cox et al., 2013; Hu et al., 2017, 2020). Besides, the Baijianshan BIF shares similar REE patterns with those BIFs from North China Craton but contains slightly higher contents of  $\text{Al}_2\text{O}_3$  and  $\text{TiO}_2$  (Li et al., 2014). Taken together, we suggest a little incorporation of terrigenous component in the Baijianshan BIF.

Positive Eu anomalies are generally interpreted as the precipitation of high-temperature hydrothermal fluids (Danielson et al., 1992; Bau and Dulski, 1999) while low-temperature hydrothermal fluids usually display weak or no Eu anomalies (Michard et al., 1993; Li et al., 2014). Most samples from the Baijianshan BIF show negligible Eu anomalies ( $\text{Eu}/\text{Eu}^* = 0.96\sim 1.14$ , averaging 1.07), which are indicative of low-temperature hydrothermal solutions (Danielson et al., 1992). The (t) value can be used as a tracer to distinguish different sources and crustal contamination. As shown in Fig. 7, the Baijianshan BIF has slightly average (t) value (average of 0.37), much lower than those of Baijianshan basalt, indicating that the deposition of the Baijianshan BIF was controlled by a hydrothermal flux with mantle- (Jacobsen and Pimentel-Klose, 1988). However, the slightly negative correlation between (Ma) and  $\text{Al}_2\text{O}_3$  contents (Fig. 8a) argues for the low proportion of detrital components inputting.

The Y/Ho ratios of BIFs can provide constraints of the material source and precipitation environment of BIFs (Hu et al., 2017). Modern seawaters have a Y/Ho ratio of 44-74 (Bau and Dulski, 1996). The terrestrial material has a Y/Ho ratio of  $\sim 26$ , any little terrestrial contamination could quickly descend seawater-like superchondritic Y/Ho ratios ( $>44$ ) (Bolhar et al., 2004). The hydrothermal fluids have almost chondritic Y/Ho ratios (26-28, Douville et al., 1999; Bau and Dulski, 1999). The Baijianshan BIF has a range of Y/Ho ratios between 23 and 32, similar to the chondritic values (26-28, Bau and Dulski, 1999), possibly due to the mixture of high and/or low-T hydrothermal fluids (Hu et al., 2020). Y/Ho ratios of the Baijianshan BIF samples might inherit from the low-T hydrothermal fluids, but the influences of terrestrial materials can not be completely excluded. Compared with the BIFs in North China Craton, the Baijianshan BIF displays relatively low positive Eu anomalies, low Y/Ho ratios as well as a slightly depleted Li et al., 2014; Wang et al., 2016), illustrating the input of low-temperature hydrothermal fluids during the deposition of the Baijianshan BIF.

Ce/Ce\*<sub>PAAS</sub> values with variable Al<sub>2</sub>O<sub>3</sub> contents of the Baijianshan BIF indicate that terrigenous detrital materials played little role in the geochemical budget of Ce. In oxidized seawater, Ce(III) transforms into Ce(IV), and then Ce(IV) is likely to be hydrolyzed and precipitated with Fe-Mn oxyhydroxides, organics and clay which lead to a significant negative Ce anomaly of seawater (Byrne and Sholkovitz, 1996). As shown in the Ce/Ce\* vs. Pr/Pr\* diagram (Fig. 10), all the Baijianshan BIF samples display no Ce anomalies, similar with most Neoproterozoic BIFs, unambiguously arguing for anoxic environment in the ancient ocean. Although the atmospheric and oceanic oxygen levels during the Neoproterozoic are still under debate, numerous studies revealed that Neoproterozoic oceans underwent a stepwise and protracted oxidation. Anoxic ferruginous deep seawater was a typical feature of the late Neoproterozoic, as inferred from geochemical proxies such as iron geochemistry (e.g., Canfield et al., 2008; Sperling et al., 2015), redox-sensitive elements (e.g., Schröder and Grotzinger, 2007; Rajabi et al., 2015), and framboidal pyrite (e.g., Rajabi et al., 2015). Stern et al. (2013) suggested that this scenario might be attributed to the seawater surfaces being covered by ice sheets, blocking the oxygen from atmosphere dissolving into the widespread anoxic ferruginous ocean. Despite no tillite was identified at the Baijianshan BIF, the coeval tillite sequences were documented in the Qurugtagh area (i.e., ca. 740 Ma Beiyixi tillite) of NE Tarim (Xu et al., 2009), which argued for the possibility of ice sheets in the Cryogenian Baijianshan ocean.

In a Fe<sup>2+</sup>-rich marine settings, both abiotic and biotic action were demonstrated as oxidation pathways for the precipitation of iron oxyhydroxides (Fe(OH)<sub>3</sub>) from dissolved Fe<sup>2+</sup> (Konhauser et al., 2011). As an element of variable valency, the fractionation of Fe isotope is affected by the oxidation-reduction condition (Bullen et al., 2001; Zhu et al., 2002; Balci et al., 2006). Experimental studies demonstrated that oxidation of Fe<sup>2+</sup> to Fe<sup>3+</sup> in solution causes considerable Fe isotope fraction, generating an enrichment of heavy Fe isotopes in Fe<sup>3+</sup> (Bullen et al., 2001; Johnson et al., 2002; Balci et al., 2006; Hou et al., 2014). On the other hand, the fractionation caused by pyrite with light-Fe isotopes is considered to play a leading role in the enrichment of heavy Fe isotopes in the residual system relative to the original state (Rouxel et al., 2016), since the pyrite is easier to precipitation from plume. Nevertheless, there is no sulfide observed in the Baijianshan BIF, thus the potential influences of sulfides on the enrichment of heavy Fe isotopes can be excluded. Under oxidized environment, the enrichment of heavy Fe isotopes is shown in oxide or hydroxide of iron trivalent (Fe<sup>3+</sup><sub>ppt</sub>) and light Fe isotope is displayed in Ferrous solution (Fe<sup>2+</sup><sub>aq</sub>) (Johnson et al., 2002). The magnitude of Fe isotope fractionation is controlled by the degree of precipitation from Fe<sup>2+</sup> to Fe<sup>3+</sup>, which is related to the degree of ocean oxidation (Yan et al., 2010). The Fe in the seawater can be completely precipitated and no fractionation of the Fe isotope occurs when the seawater is completely oxidized, thus the Fe isotope values in iron oxide precipitates (Fe<sup>3+</sup>) can represent the Fe isotope information of seawater. However, when the seawater is partially oxidized, the Fe in the seawater is partially precipitated and the Fe isotope fractionation will occur in iron oxide precipitates (Fe<sup>3+</sup>). Therefore, the Fe isotope value can be used as a proxy to decipher the redox state of seawater (Yan et al., 2010). Compared with some BIFs deposited during the Archean-Paleoproterozoic (Dauphas et al., 2004; Rouxel et al., 2005; Li et al., 2012; Johnson et al., 2003; Hou et al., 2014; Li et al., 2014), the Neoproterozoic Baijianshan Sanjiang (Yan et al., 2010) and Xinyu BIFs (Shen et al., 2008) have higher values <sup>57</sup>Fe (Fig. 11). Several studies demonstrated that the low <sup>57</sup>Fe values of BIFs could be due to the contributions from a continental component and the continental Fe source is best explained by Fe mobilization on the continental margin by microbial dissimilatory iron reduction (DIR) (Johnson et al., 2008; Li et al., 2015). As for the Fe isotope values of Archean-Paleoproterozoic BIFs we cited, the Fe source of SW Greenland BIF, Anshan-Benxi BIFs, Gongchangling BIFs and Yuanjiacun BIFs are unambiguously demonstrated to be hydrothermal type similar to the scenario of the Baijianshan BIF (Dauphas et al., 2004; Li et al., 2012; Li et al., 2014; Hou et al., 2014), while the Fe of Transvaal, Manjeri, Belingwe Belt and Zimbabwe BIFs are probably from multiple sources (abiologic and biologic processes) (Johnson et al., 2008; Grassineau et al., 2011). The comparison reveals that the environment in which they precipitated was more reduced. Both the iron isotope and elemental geochemistry thus demonstrate that the Neoproterozoic Baijianshan ocean might have the same reduced environment as Archean or Paleoproterozoic, or even more reduced.

### 6.3 A model of the Baijianshan BIF

During the late Neoproterozoic period (i.e., 750-550 Ma), the Earth experienced long-lived global glaciations, known as the Snowball Earth events (Hoffman et al., 1998). In view of the Neoproterozoic oxygenation event (NOE) at that time, the content of oxygen in atmosphere was much higher than before (Och et al., 2012). Recently, El-Rahman et al. (2020) argued that some Neoproterozoic BIFs are not glaciogenic due to their significant depleted Nd isotope compositions. Under this scenario, the iron was most likely deriving from hydrothermal alteration of juvenile oceanic crust. The Chondritic-like Nd isotope compositions of the Baijianshan BIF ( do not favor this model. Furthermore, the close temporally-spatially related tillite sequences in the Qurugtagh of NE Tarim (Xu et al., 2009), 765 Ma-735 Ma tillites in Namibia, suggested a local glaciation termed as Kaigas

glaciation (Kay et al., 2001; Cailteux et al., 2005; Hoffmann et al., 2006). Considering the high latitude position of the Tarim in the Rodinia configuration (Li et al., 2008), the Kaigas glaciation (corresponding to the Beiyixi tillite in Tarim) most possibly occurred in Tarim. In combination with previous studies on other BIFs with the data presented in this study, we construct a cartoon model showing the formation mechanism of the Baijianshan BIF (Fig. 12).

Stage 1, the seawater was covered by ice sheet which led to the insulation between atmosphere and oceans with the in

And then it induced the emergence of a reductive ocean. The iron supplied by low-temperature hydrothermal fluids could exist as the ferrous form, forming a relatively stable anoxic ferruginous reservoir in the ocean.

Stage 2, at interglacial period, the partial melting of the ice sheet caused some oxygen from the atmosphere into the ocean, which emerged a stratified seawater. There were mainly anoxic and Fe<sup>2+</sup>-rich deeper seawater, whereas the Fe<sup>2+</sup> was gradually oxidized into Fe<sup>3+</sup> and precipitated into iron oxides in near-shore oxic shallow seawaters. The low-temperature hydrothermal was the dominant source of iron for the precipitation process to interpret the slight positive Eu anomalies and positive values of  $\delta^{56}\text{Fe}$ . Under the effect of the melting ice-sheet, the increasing weathering then contributed to a large amount of terrestrial materials into the shallow seawater, resulting the precipitation of the iron-rich BIF.

## Conclusions

The Baijianshan BIF precipitated in glacially-influenced settings during the Neoproterozoic (737 Ma). Low-T hydrothermal fluids, accompanied with limited detritus input, were the dominated iron sources for the formation of Baijianshan BIF.

The Baijianshan ocean was initially covered by ice sheets, giving rise to the extremely reduced environment favored by Fe<sup>2+</sup>. Then it was oxidized into Fe<sup>3+</sup> and precipitated as magnetite due to the thawing of ice sheets. At the early stage, local reduced ocean enhanced Fe fluxes genetically related to the intense magmatism and hydrothermal activities due to the existence of ice sheets. Thawing induced oxidized oceanic layer, favoring the formation of the Neoproterozoic BIF at the Baijianshan.

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Figure captions

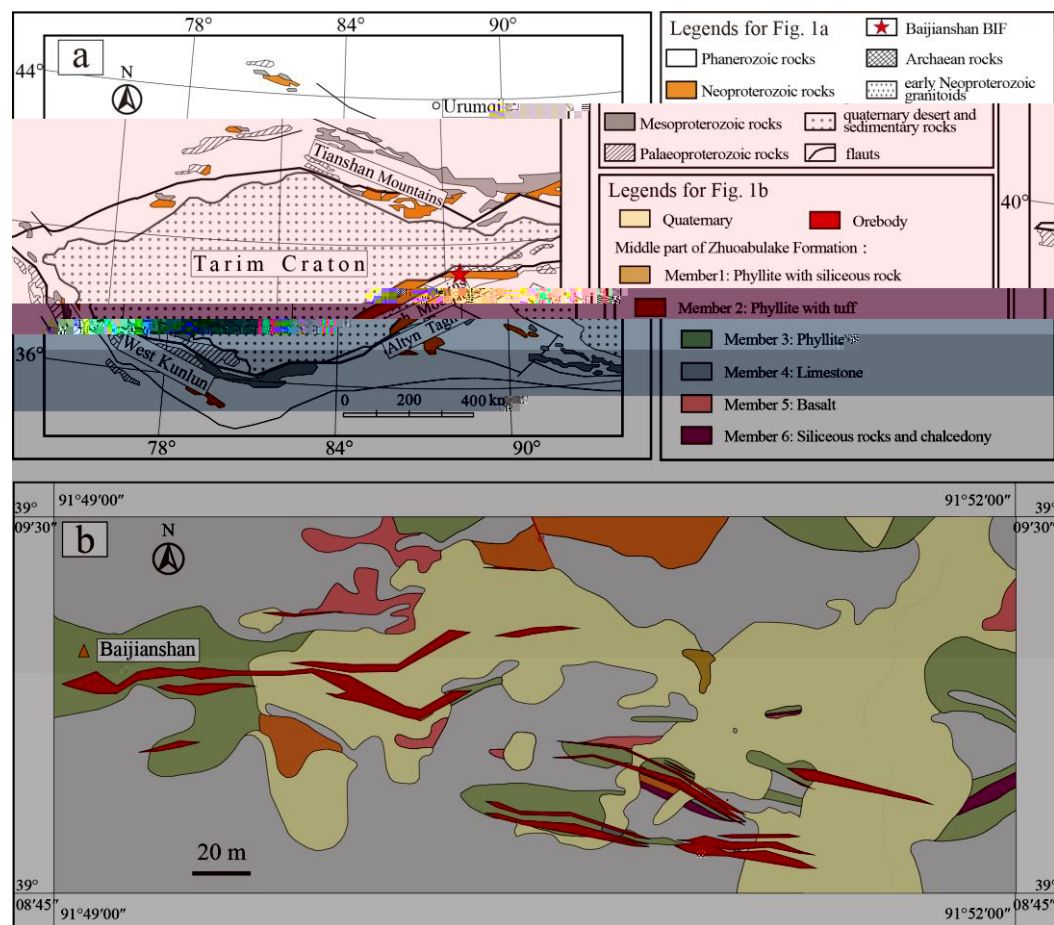


Fig. 1 (a) Sketch Precambrian geological map of Tarim Craton, showing the location of the Baijianshan BIF (modified after Lu et al., 2008). (b) Geological map of the Baijianshan Banded Iron Formation (Baijianshan BIF).

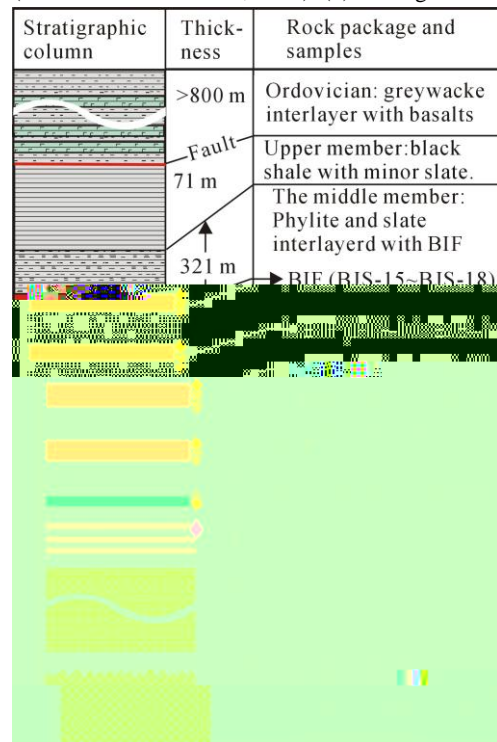


Fig. 2 Stratigraphical column of the Baijianshan BIF.

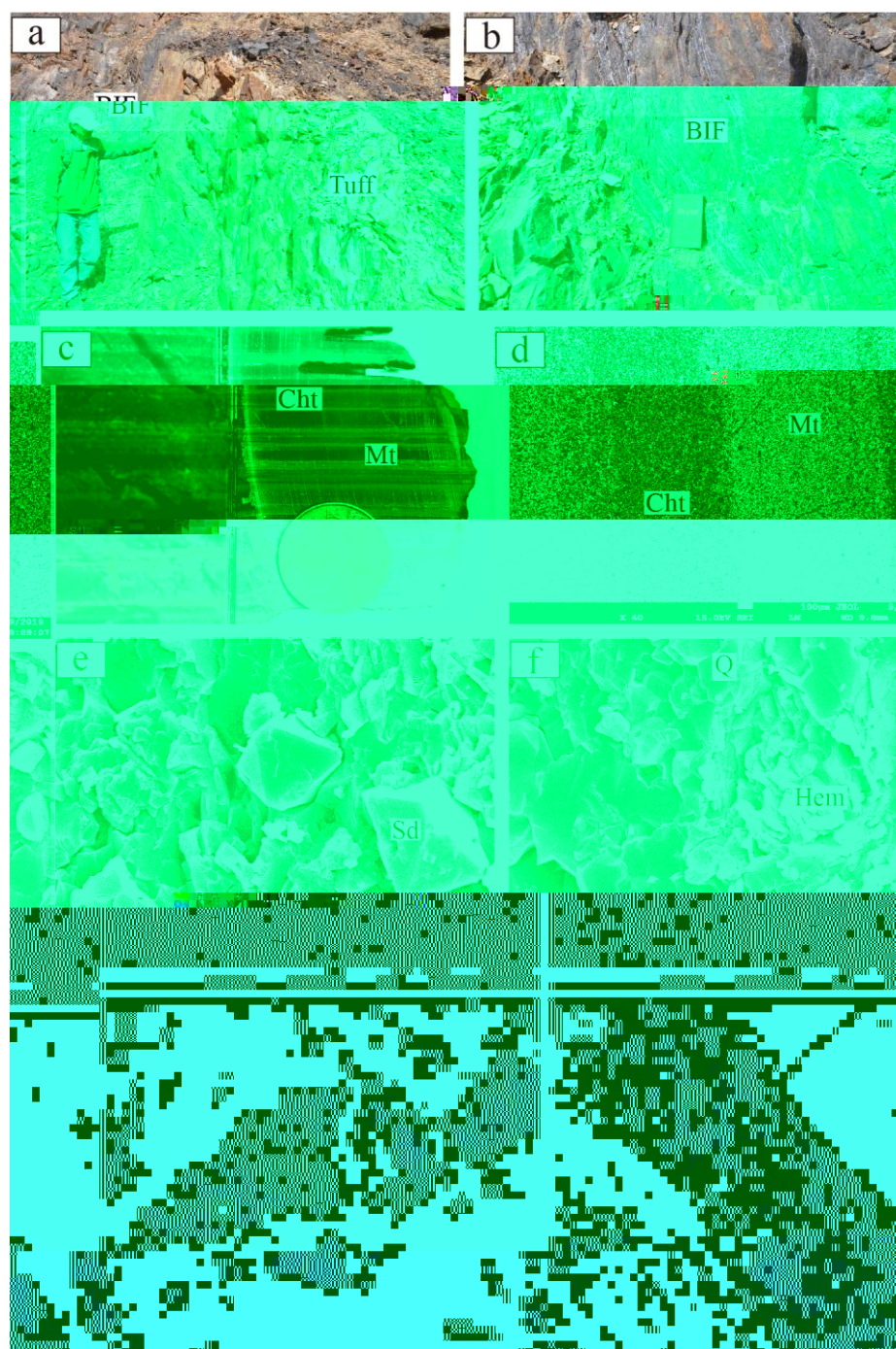


Fig. 3 Field photographs, hand specimen and mineral features of the Baijianshan BIF. (a) The tuff layers in the Baijianshan BIF. (b) Laminated Fe orebodies with local bend. (c) hand specimen. (d) Scanning electron microscopy (SEM) images. Detailed petrographic examination by SEM and transmitted light showing mineral assemblage: (e) Siderite. (f) Hematite. (g) Chlorite. (h) Sericite. Cht = Chert, Mt = Magnetite, Sd = Siderite, Q = Quartz, Hem = Hematite, Chl = Chlorite, Ser = Sericite.

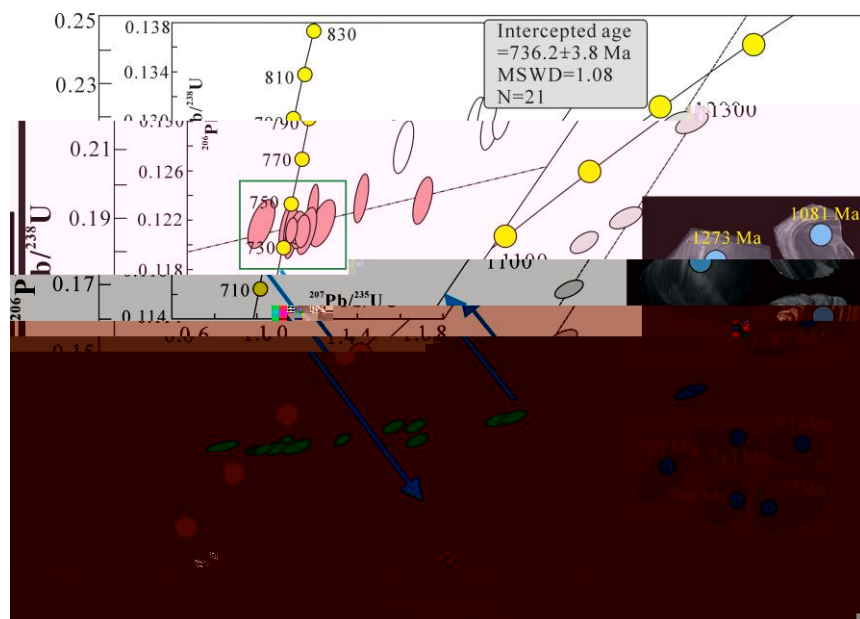


Fig. 4 Concordia diagram of the zircon U-Pb ages of the tuff at the Baijianshan BIF.

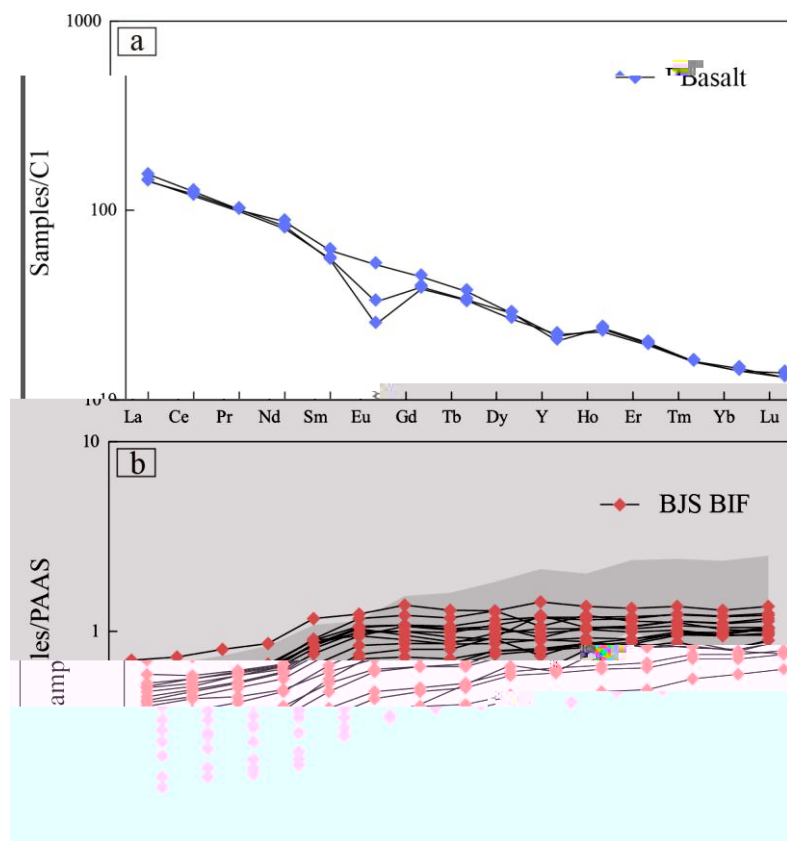


Fig. 5 (a) Chondrite-normalized REE pattern of the basalts. Chondrite-normalized values after Sun and McDonough (1989). (b) PAAS-normalized REE+Y diagrams of BIFs from the Baijianshan BIF. The grayish zone represents the REY data from the Xinyu BIF (Li et al., 2014). PAAS-normalized values after McLennan (1989).

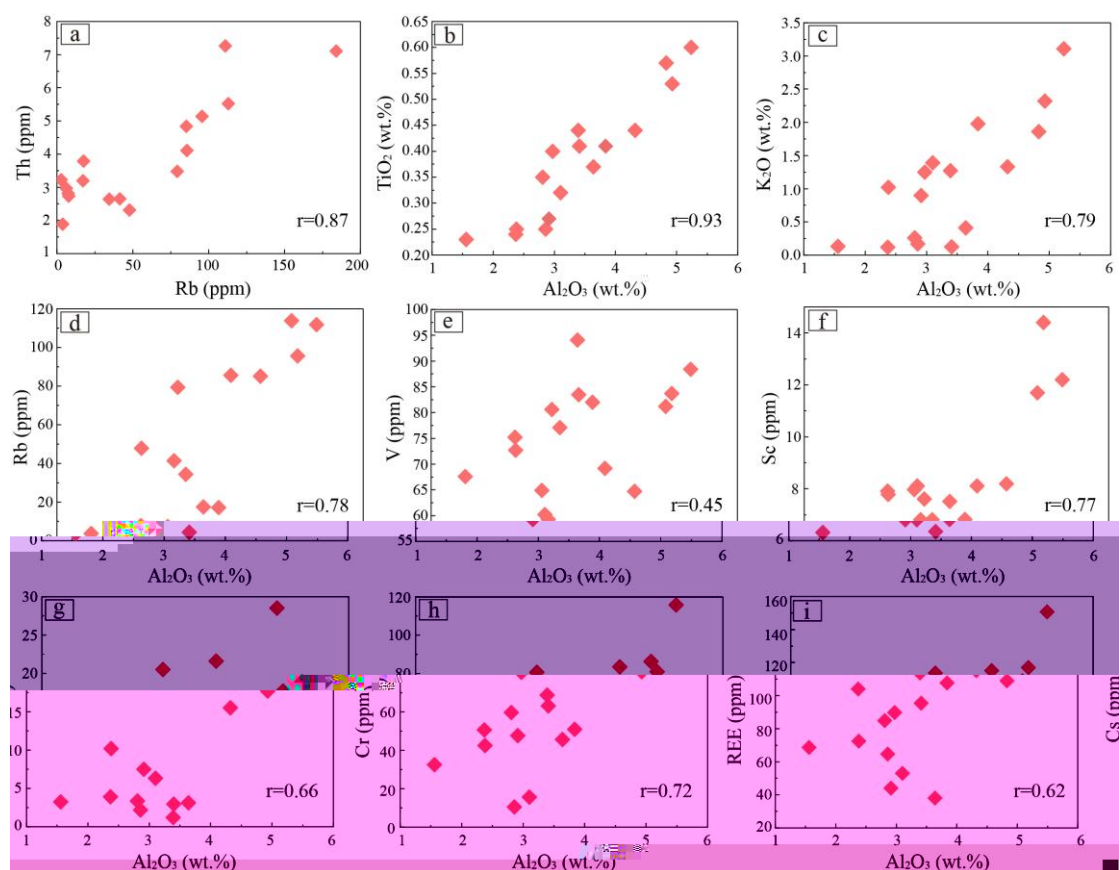
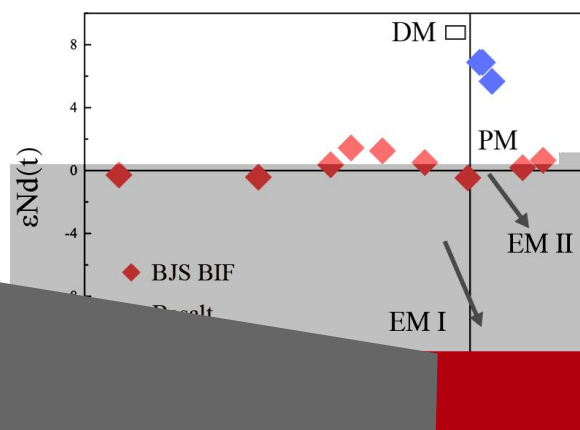


Fig. 6 (a) Bi-variation between Rb and Th of Baijianshan BIF samples; (b-i) Bi-variation between  $\text{Al}_2\text{O}_3$  and  $\text{TiO}_2$  (b),  $\text{K}_2\text{O}$  (c), Rb (d), V (e), Sc (f), Cs (g), Cr (h) and REE (i) of the Baijianshan BIF. Data for bulk-rock BIF samples are listed in supplementary Table 2.



and basalt samples in Baijianshan.



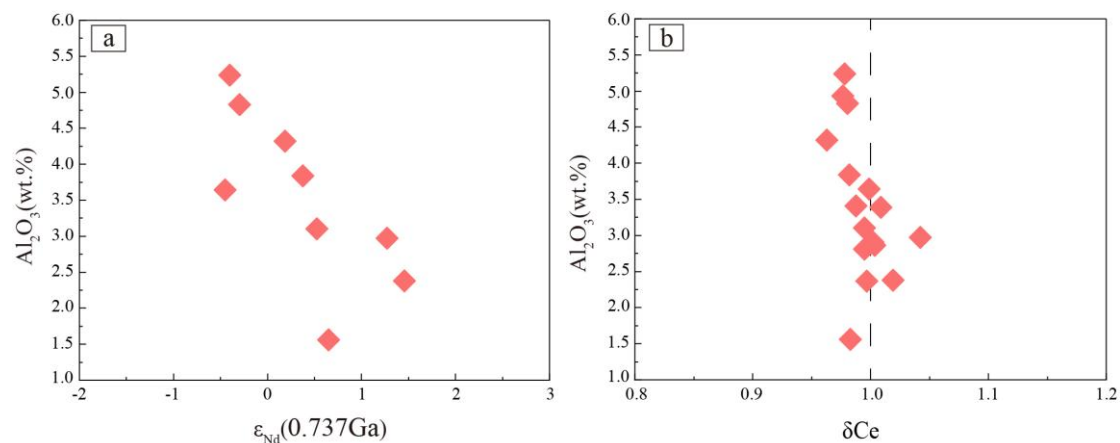


Fig. 8 (a)  $\epsilon_{\text{Nd}}(0.737\text{Ga})$  vs.  $\text{Al}_2\text{O}_3$  diagram of the Baijianshan BIF; (b)  $\text{Ce}/\text{Ce}^*_{\text{PAAS}}$  vs.  $\text{Al}_2\text{O}_3$  diagram of the Baijianshan BIF.

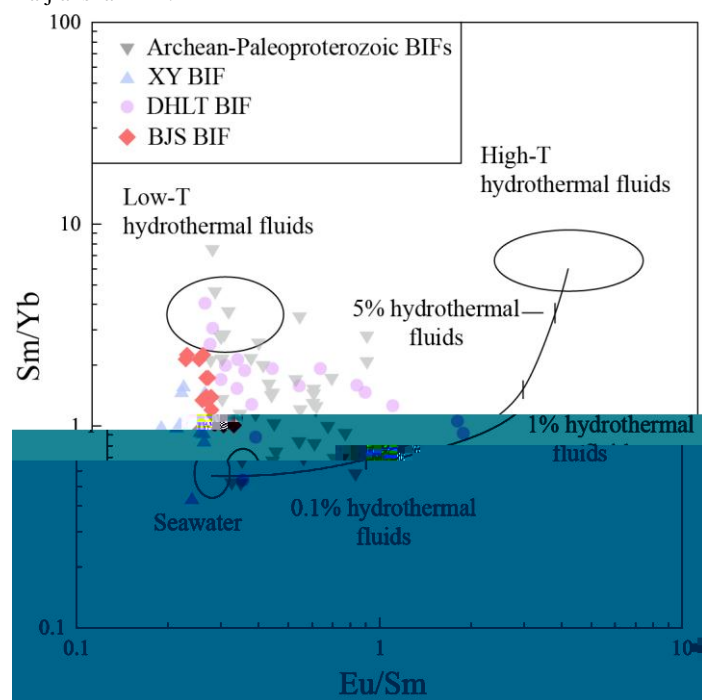


Fig. 9 Sm/Yb vs. Eu/Sm ratios illustrating two component mixing model of the hydrothermal fluid of the Baijianshan BIF (after Alexander et al., 2008). BJS-Baijianshan BIF; XY-Xinyu BIF (Li et al., 2014); DHLT-Dahongliutan BIF (Hu et al., 2017). Data sets of Archean-Paleoproterozoic BIFs are from Planavsky et al. (2010). Average compositions of high-T ( $> 300\text{ }^{\circ}\text{C}$ ) hydrothermal fluids, low-T ( $< 200\text{ }^{\circ}\text{C}$ ) hydrothermal solutions and Pacific seawaters were after Bau and Dulski (1999), Michard et al. (1983) and Alibo and Nozaki (1999), respectively.

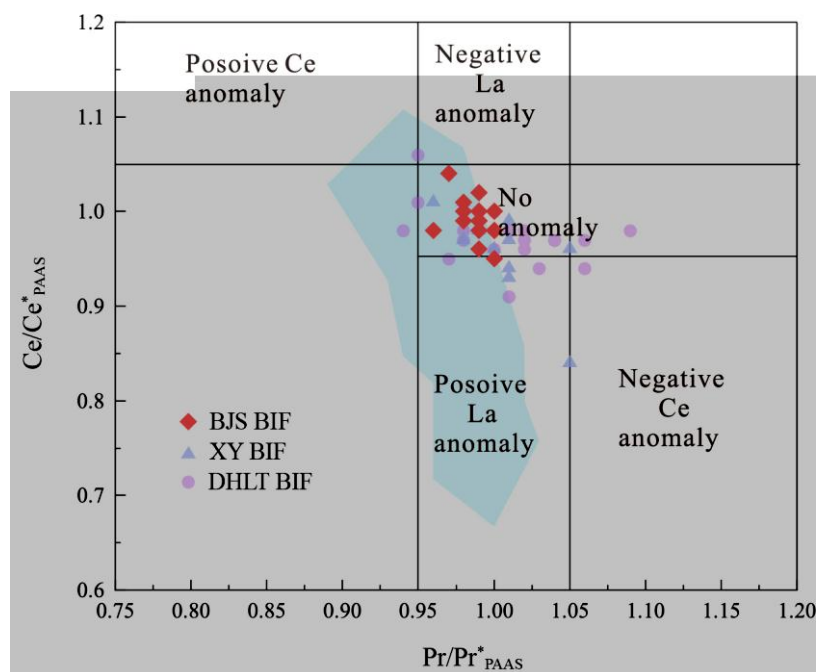


Fig. 10  $\text{Pr}/\text{Pr}^*_{\text{PAAS}}$  vs.  $\text{Ce}/\text{Ce}^*_{\text{PAAS}}$  diagram for the BIF at Baijianshan (after Bau and Dulski, 1996). The light blue area represents the data of Archean-Paleoproterozoic BIFs. Data sets of other BIFs are same as those of Fig. 9.

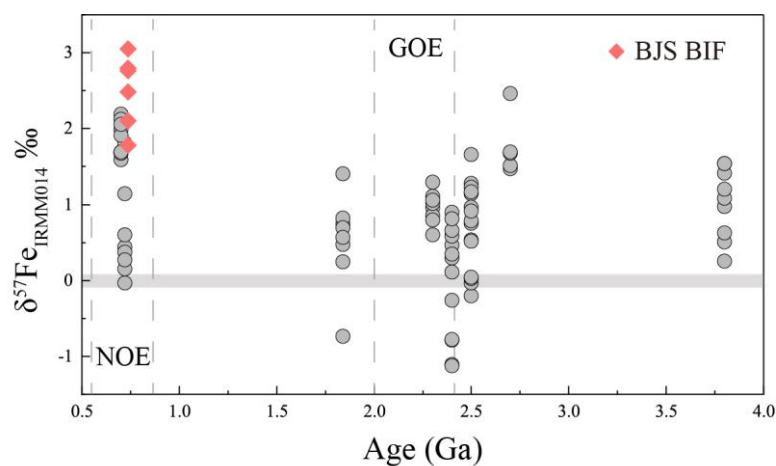


Fig. 11  $\delta^{57}\text{Fe}_{\text{IRMM-014}}$  of various ages BIFs in the world (Data of SW Greenland BIFs are from Dauphas et al.(2004); Manjeri Formation, Belingwe Belt, Zimbabwe BIFs are from Rouxel et al. (2005); Anshan-Benxi BIFs are from Li et al. (2012); Transvaal BIFs are from Johnson et al. (2003); Yuanjiacun BIFs are from Hou et al. (2014); Gongchangling BIFs from Li et al. (2014); Sanjiang BIF are from Yan et al. (2010) and Xinyu BIF are from Shen et al. (2008)).



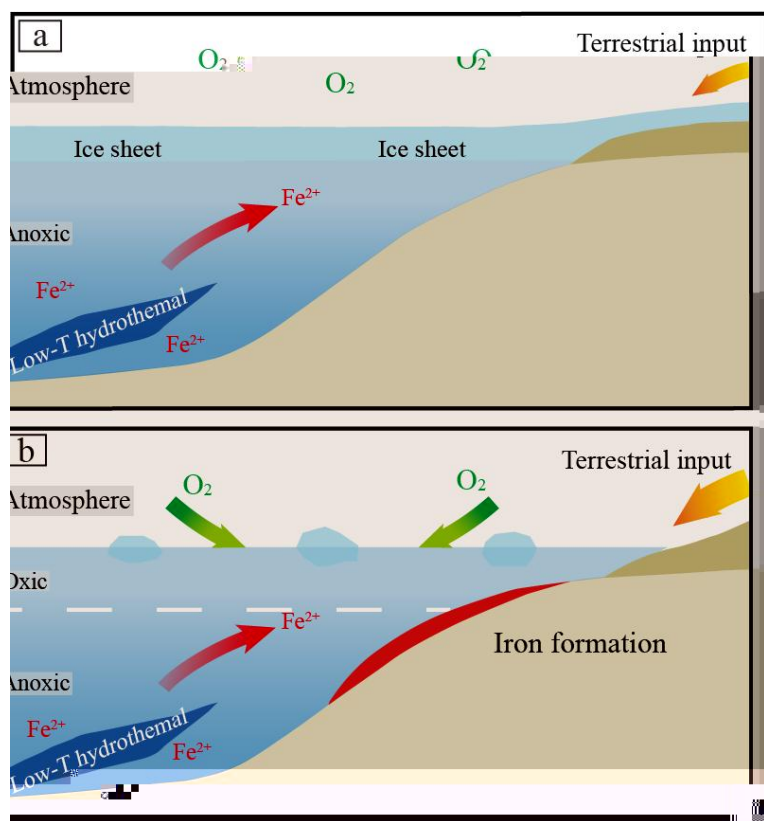


Fig. 12 Cartoon model of the formation of the Baijianshan BIF (see details in the text).

## Appendix Table Captions

Appendix-Table 1 LA-ICP-MS U-Pb data for zircon from tuff in the first section of the Lapeiquan Formation

Appendix-Table 2 Geochemical compositions of the Baijianshan BIF

Appendix-Table 3 Sr-Nd isotopic data of the Baijianshan BIF

Appendix-Table 4 Fe isotopic data of the Baijianshan BIF

## Supplementary Table 1 Zircon U-Pb age data of the tuff from the Baijianshan BIF

Spot	U ppm	Th ppm	Th/U	<sup>206</sup> Pb/ <sup>238</sup> UAge		<sup>207</sup> Pb/ <sup>235</sup> UAge		<sup>207</sup> Pb / <sup>206</sup> Pb	<sup>207</sup> Pb/ <sup>235</sup> U	<sup>206</sup> Pb/ <sup>238</sup> U	<sup>206</sup> Pb/ <sup>238</sup> U	<sup>206</sup> Pb/ <sup>238</sup> U	
2071-1	3079	831	0.27	1335	14	1360	21	0.0887	0.0012	2.8156	0.0427	0.2301	0.0025
2071-2	805	427	0.53	898	11	1154	20	0.1027	0.0014	2.1169	0.0369	0.1495	0.0018
2071-3	529	381	0.72	776	8	1009	18	0.0965	0.0016	1.7019	0.0307	0.1279	0.0014
2071-4	155	81	0.52	835	9	1350	22	0.1458	0.0022	2.778	0.0462	0.1382	0.0015
2071-5	184	101	0.55	751	8	1009	18	0.0999	0.0016	1.7022	0.0304	0.1236	0.0013
2071-6	297	252	0.85	1273	14								

**Supplementary Table 2 Geochemical compositions of Baijianshan BIF**

Sample	BJS-1	BJS-2	BJS-3	BJS-4	BJS-5	BJS-6	BJS-7	BJS-8	BJS-9	BJS-10
Major elements (%)										
Fe <sub>2</sub> O <sub>3</sub>	53.60	41.52	56.64	52.26	30.01	36.85	41.01	45.22	39.78	48.55
SiO <sub>2</sub>	29.91	38.96	26.81	31.20	40.92	40.79	36.97	41.77	43.66	37.59
Al <sub>2</sub> O <sub>3</sub>	1.56									

TFe <sub>2</sub> O <sub>3</sub>	21.79	28.86	44.82	42.63	26.98	13.11	40.45	11.61	11.70	11.98
SiO <sub>2</sub>	55.06	49.87	41.24	43.65	53.64	59.90	42.80	39.94	40.91	44.55
Al <sub>2</sub> O <sub>3</sub>	5.24	4.93	2.81	2.86	4.83	7.58	3.84	11.36	11.45	10.76
CaO	0.81	1.02	0.75	0.75	0.76	0.84	0.65	9.77	9.31	7.45
MgO	1.57	1.51	0.74	0.86	1.36	2.32	1.31	8.40	8.54	10.38
TiO <sub>2</sub>	0.60	0.53	0.35	0.25	0.59	0.86	0.41	3.66	3.80	3.60
MnO	0.03	0.04	0.05	0.03	0.03	0.02	0.03	0.24	0.25	0.23
P <sub>2</sub> O <sub>5</sub>	0.20	0.35	0.49	0.40	0.23	0.19	0.54	0.46	0.43	0.38
K <sub>2</sub> O	3.11	2.32	0.26	0.17	1.86	3.78	1.98	1.35	1.27	0.36
Na <sub>2</sub> O	0.94	1.25	0.14	0.19	0.63	0.77	0.88	0.94	0.98	0.59
LOI	2.06	1.97	4.57	2.89	2.9	3.19	1.92	12.02	11.39	9.60
FeO	7.97	7.26	3.89	5.29	6.06	6.57	5.31	-	-	-
Total	99.38	99.91	100.11	99.97	99.87	99.13	100.12	99.75	100.03	99.88
Trace elements (ppm)										
Cr	118	81.0	59.7	10.5	86.2	112	50.9	214	214	306
Ni	15.6	14.5	12.1	11.0	15.6	12.3	12.6	162	166	217
Co	76.1	70.2	33.7	34.7	54.0	27.1	45.3	48.0	62.9	50.7
Sc	12.2	14.4	7.97	8.11	11.7	15.1	8.11	26.4	26.1	25.7
Ba	662	562	111	19.2	359	1219	385	877	864	296
Sr	45.0	53.3	45.8	26.1	32.6	37.0	48.9	205	209	150
Rb	111	95.6	7.36	6.13	113	184	85.6	34.4	32.2	12.3
Zr	101	93.4	60.0	87.9	79.9	116	79.7	259	251	228
Hf	2.63	2.37	1.23	1.50	2.11	2.99	1.92	5.35	5.74	6.00
Ta	0.88	0.75	0.37	0.36	0.64	0.84	0.53	2.24	2.33	2.14
Ga	7.95	7.63	4.39	4.63	7.12	10.3	5.97	18.3	17.2	18.8
Cu	27.6	21.6	8.12	3.56	21.4	36.0	12.4	-	-	-
Zn	62.1	40.0	26.0	23.8	35.3	37.0	29.6	-	-	-
Pb	7.80	8.24	3.54	5.32	7.64	15.4	5.75	-	-	-
U	0.47	0.69	0.39	0.40	0.72	0.82	0.37	1.45	1.09	0.92
Th	7.26	5.14	2.81	2.97	5.52	7.11	4.11	3.32	3.37	3.18
Nb	8.35	8.21	4.19	4.86	6.80	9.50	6.11	38.0	36.5	33.6
La	26.8	20.3	13.0	9.82	19.1	22.5	17.3	36.9	34.1	33.7
Ce	58.3	44.2	30.2	22.8	42.0	46.1	39.0	77.2	72.9	74.9
Pr	7.03	5.35	3.73	2.77	5.09	5.51	4.81	9.37	9.15	9.28
Nd	29.3	22.4	16.2	11.7	21.3	22.6	20.7	37.9	36.6	40.0
Sm	6.46	4.95	3.67	2.65	4.70	4.96	4.70	8.06	8.20	9.11
Eu	1.33	1.15	0.91	0.66	1.06	1.01	1.12	1.41	1.85	2.91
Gd	6.39	5.03	4.02	3.03	4.64	4.93	5.01	7.69	7.82	8.87
Tb	1.00	0.80	0.67	0.52	0.72	0.79	0.82	1.21	1.22	1.35
Dy	5.98	4.98	4.53	3.71	4.26	4.84	5.36	6.57	7.04	7.04
Y	28.5	24.9	27.9	22.0	20.1	26.1	31.9	34.8	34.0	32.1
Ho	1.20	1.04	1.02	0.85	0.86	1.01	1.17	1.25	1.29	1.31
Er	3.23	2.94	3.00	2.59	2.33	2.82	3.38	3.11	3.15	3.20
Tm	0.45	0.43	0.46	0.40	0.34	0.41	0.50	0.39	0.40	0.39
Yb	2.87	2.93	3.06	2.76	2.20	2.69	3.36	2.28	2.36	2.28
Lu	0.43	0.45	0.49	0.45	0.33	0.41	0.53	0.34	0.32	0.32
	150.77	116.95	84.95	64.71	108.93	120.58	107.75	193.68	186.4	194.67
Y/Ho	23.75	23.94	27.35	25.76	23.29	25.84	27.26	27.84	26.36	24.50
La/La*	1.06	1.07	1.09	1.02	1.06	1.09	1.10	1.14	1.06	1.14
Ce/Ce*	0.98	0.98	0.99	1.00	0.98	0.95	0.98	0.98	0.98	1.01
Y/Y*	0.85	0.88	1.04	0.99	0.84	0.95	1.02	0.89	0.83	0.78
Pr/Pr*	1.00	1.00	0.99	0.99	1.00	1.00	0.99	0.97	0.99	0.95
Eu/Eu*	0.97	1.08	1.10	1.08	1.07	0.96	1.08	0.54	0.69	0.97
La <sub>N</sub> /Yb <sub>N</sub>	0.69	0.51	0.31	0.26	0.64	0.62	0.38	10.99	9.82	10.04

Supplementary Table 3 Sr-Nd isotopic data of the Baijianshan BIF

Sample	<sup>87</sup> Rb/ <sup>86</sup> Sr	<sup>87</sup> Sr/ <sup>86</sup> Sr	( <sup>87</sup> Sr/ <sup>86</sup> Sr) <sub>i</sub>	<sup>147</sup> Sm/ <sup>144</sup> Nd	<sup>143</sup> Nd/ <sup>144</sup> Nd	T <sub>DM</sub> (Ga)		
BJS-1	0.0433	0.717927	0.000004	0.7175	0.512425	0.000006	1.63	0.65
BJS-4	4.2608	0.727751	0.000009	0.6829	0.512394	0.000003	1.39	1.46
BJS-5	1.4305	0.728893	0.000008	0.7138	0.512363	0.000003	1.58	0.19
BJS-7	2.4192	0.729423	0.000009	0.7040	0.512376	0.000002	1.78	-0.45
BJS-8	2.8555	0.718681	0.000003	0.6886	0.512351	0.000004	1.36	1.27
BJS-10	2.6390	0.723954	0.000005	0.6962	0.512364	0.000002	1.51	0.53
BJS-12	7.0891	0.740795	0.000007	0.6662	0.512311	0.000003	1.59	-0.40
BJS-16	9.9936	0.746329	0.000008	0.6412	0.512317	0.000003	1.58	-0.30
BJS-18	5.0054	0.731954	0.000008	0.6793	0.51237	0.000004	1.56	0.38
2702H1	0.4737	0.711034	0.000009	0.7061	0.512662	0.000003	0.88	6.90
2702H2	0.4355	0.711156	0.000009	0.7066	0.512694	0.000006	0.89	6.88
2702H3	0.2322	0.710822	0.000012	0.7084	0.512644	0.000005	1.02	5.69

Supplementary Table 4 Fe isotopic data of the Baijianshan BIF

Sample	<sup>57</sup> Fe <sub>IRMM-014</sub>	<sup>56</sup> Fe <sub>IRMM-014</sub>
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BJS-2	2.10	0.09	1.46	1.46
BJS-5	1.78	0.02	1.24	1.24
BJS-7	2.76	0.09	1.89	1.89
BJS-9	2.49	0.05	1.69	1.69
BJS-12	3.05	0.04	2.17	2.17
BJS-17	2.80	0.01	1.91	1.91

Ac